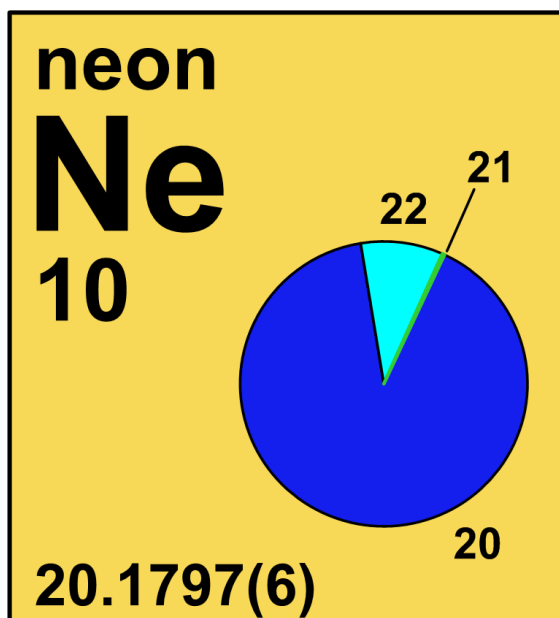





## 4.10 neon



Stable isotope	Relative atomic mass	Mole fraction
$^{20}\text{Ne}$	19.992 440 18	0.9048
$^{21}\text{Ne}$	20.993 8467	0.0027
$^{22}\text{Ne}$	21.991 3851	0.0925

## Half-life of radioactive isotope

Less than 1 hour   
 Between 1 hour and 1 year   
 Greater than 1 year 



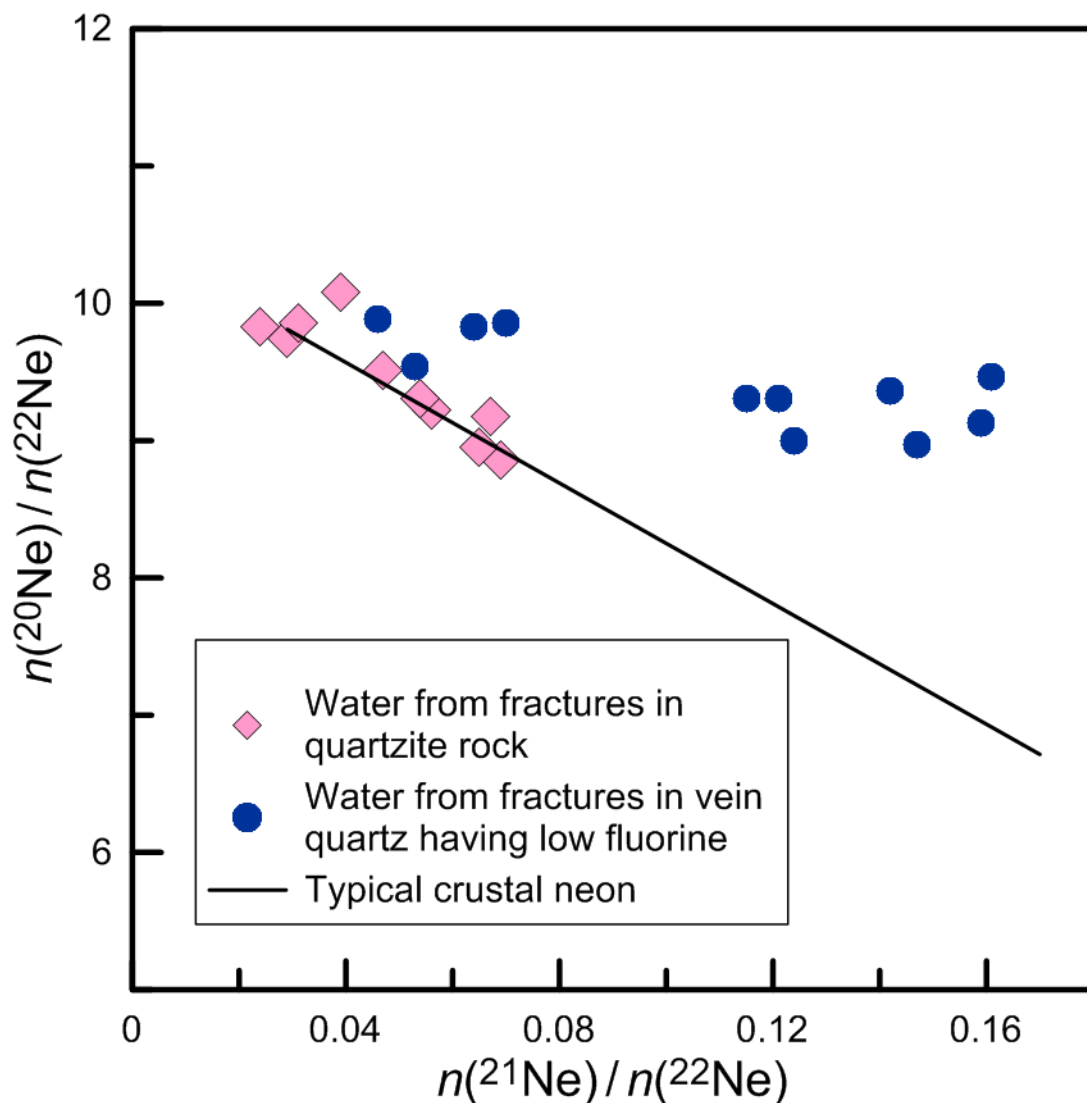
## 4.10.1 Neon isotopes in Earth/planetary science

Neon is subject to stable **isotopic fractionation** by physical processes such as exchange between gas, liquid, and solid phases. Small variations in the **isotope-amount ratio**  $n(^{22}\text{Ne})/n(^{20}\text{Ne})$  have been used to examine gas-liquid exchange processes during groundwater recharge (water moving downward from the surface) and discharge [26, 98, 99].

## 4.10.2 Neon isotopes in geochronology

Some  $^{21}\text{Ne}$  and  $^{22}\text{Ne}$  form naturally in the Earth's crust largely by reactions of  $^{18}\text{O}$  and  $^{19}\text{F}$  in minerals with **neutrons** and **alpha particles** emitted from uranium and thorium decay, called nucleogenic neon isotopes [26, 98]. In addition, neon **isotopes** can form at the surface of the Earth and in extraterrestrial bodies by cosmic-ray-induced **spallation** reactions on magnesium, silicon, aluminum, and sodium [100, 101]. Analyses of all three stable neon isotopes may be used to distinguish these sources from **primordial** neon. The relative amounts of atmospheric neon and crustal nucleogenic neon isotopes in deep groundwaters and natural gases have been

used in studies of solid-water-gas interactions and migration (Figure 4.10.1). The **cosmogenic** component is mainly detected in  $^{21}\text{Ne}$  and can be used to determine cosmic-ray exposure ages of rock samples, including **meteorites** exposed during travel through space and boulders exposed by melting of glacial ice (Figure 4.10.1).



**Fig. 4.10.1:** Neon-isotope ratios of water from fractures in quartzite (open diamonds) and water from fractures in vein quartz (solid circles) from the deep gold mines of the Witwatersrand Basin, South Africa [102]. The isotope-amount ratio  $n(^{21}\text{Ne})/n(^{22}\text{Ne})$  depends upon the mole ratio of oxygen to fluorine in the  $\sim 40\text{-}\mu\text{m}$  reaction range of **alpha particles** from uranium and thorium. Lippmann-Pipke *et al.* [102] show that the neon end-member represents a fluorine-depleted fluid component that was trapped in fluid inclusions in vein quartz more than  $2 \times 10^9$  years ago.

## IUPAC

### 4.10.3 Neon isotopes in industry

Masers (Microwave Amplification by Stimulated Emission of Radiation) containing  $^{20}\text{Ne}$  have been used to study quantum physics.  $^{21}\text{Ne}$  may also play a role in maser studies of quantum physics [103].

### 4.10.4 Neon isotopes used as a source of radioactive isotope(s)

$^{22}\text{Ne}$  is used to produce the **radioisotope**  $^{22}\text{Na}$  via the reaction  $^{22}\text{Ne}(\text{p}, \text{n})^{22}\text{Na}$  [104].  $^{20}\text{Ne}$  has been used to produce the radioisotope  $^{18}\text{F}$  via the reaction  $^{20}\text{Ne}(\text{d}, ^4\text{He})^{18}\text{F}$  [104].